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Abstract

The variety and wideness of substitutions shown by tourmaline group, mainly in the Z, Y and X sites, make possible a high number of real and hypothetical end-members. However in some cases it is not so obvious to determine to what end-members to refer to, but the wide stability field of tourmaline and its occurrence in different geological environments make useful to define the composition as percentages of end-members to trace back to its genesis. In this paper the followed procedure to identify the components of a complex tourmaline is given. During a study on the granitoid intrusions outcropping on Sithonia, Chalkidiki Peninsula (Northern Greece), a swarm of very small crystals of tourmaline, subhedral or more often allotriomorphic, has been found as stuffing of subparallel micro fractures in the quartz core of a pegmatite dike. Their composition, analysed by EMPA, turns out to be rather unusual and not satisfactorily referable to so far proposed solid solutions. It is characterized by remarkable Al deficiency (ZAl < 6 apfu), high amounts of magnesium and iron with Mg > Fe and Na dominating X site. On the basis of compositional and crystal-chemical evidences, supported by previous literature, it was hypothesized that Mg and Fe have to be referred to an uvite–feruvite component, while the considerable part of total iron in its trivalent oxidation state yields a povondraite component according also with Na in X site. These data, coupled with a theoretical evaluation of bond valence sum (BVS) that allowed inferring the presence of OH anions in the W site, make reasonable to consider Sithonia tourmaline as the new variety sodian povondraite hydroxil uvite-feruvite. On the basis of its composition, of its fabric and of the oscillatory trends shown by the main elements of Y and X sites, this tourmaline may be considered post magmatic likely of hydrothermal genesis.

Key words: crystal-chemistry, end-members, uvite-feruvite, povondraite, pegmatite.

Περίληψη

Η ποικιλία και το εύρος των υποκαταστάσεων που παρουσιάζονται στις θέσεις Z, Y και X, καθώς και η παρουσία του προσδιορισμού ενός σημαντικού αριθμού πραγματικών και υποθετικών ακραίων μελών. Δεν είναι δυνατόν να ορίσουμε ακραίο μέλος, αλλά η ποικιλία και η διάσταση του τουρμαλίνη, συγκεκριμένα στις θέσεις Z, Y και X, καθώς και η υποκατάσταση Mg > Fe και Na στη θέση X, δίνουν περισσότερη πληροφορία για τη συστατική του τουρμαλίνη. Στην βάση των συστατικών και κρυσταλλοχημικών δεδομένων και με βάση την προηγούμενη λίτη, καταχωρήθηκε ότι οι Mg και Fe πρέπει να αντιστοιχούν σε έναν ουβίτη–φερουβίτη και η σημαντική ποσότητα του συνολικού ισχίου από αριθμό αποκαλύφθηκε ως ποβόνδραίτη και με βάση την προηγούμενη λίτη, καταχωρήθηκε ότι η συστατική του τουρμαλίνη θα πρέπει να αντιστοιχούν σε έναν ποβόνδραίτη. Ως βασικά μέταλλα της θέσης Υ και Χ, η ταξινόμηση της συστατικής του τουρμαλίνη μπορεί να αναγνωριστεί ως προσδιορισμός του υδροθερμικού γενεσίου.
ζήτηση αυτό δεν είναι τόσο εύκολο. Ωστόσο, ο προσδιορισμός της χημικής σύστασης του βοηθά στον καθορισμό παραμέτρων που σχετίζονται με το περιβάλλον γέννησή του. Σε αυτή την εργασία παρατίθεται η μεθοδολογία που ακολούθησε για τον προσδιορισμό των συστατικών μιας σύνθετης μορφής τουρμαλίνη που απαντά στη χερσόνησο της Σιθωνίας. Ο τουρμαλίνης αυτός, που έχει υπόλοιπο ή πιο συχνά αλλοτριόμορφο σχήμα, εντοπίστηκε να γεμίζει με πλήθος μικρών κρυστάλλων του παράλληλες μικρο-διακλάσεις στον χαλαζιακό πυρήνα πηγματιτικής φλέβας. Η σύσταση του οδηγεί στο συμπέρασμα ότι πρόκειται μάλλον για μέλος ασυνήθους στερεού διαλύματος. Χαρακτηρίζεται από σημαντικό έλλειμμα ΑΙ ( ΑΙ < 6 apfu), υψηλές ποσότητες μαγνησίου και σιδήρου με Mg > Fe και Na που υπερτερεί στη θέση Χ. Με βάση τη σύσταση και το χημισμό του κρυστάλλου, θεωρείται ότι το Mg και ο Fe σχετίζονται με μέλος της σειράς ουβίτη - φερουβίτη, ενώ ένα σημαντικό μέρος του ολικού σιδήρου σε τρισθενή μορφή δίνει ένα ποβονδραϊτικό συστατικό, σύμφωνα επίσης με το θεωρητικό υπολογισμό του αθροίσματος στη θέση W, μας ενισχύει την άποψη ότι ο τουρμαλίνης της Σιθωνίας ανήκει σε νέα ποικιλία με ονομασία νατρούχος ποβονδραϊτικός υδροξυλ- ουβίτης-φερουβίτης. Με βάση τη σύσταση, τη δομή και τις τάσεις ζώνωσης που παρουσιάζουν τα κύρια στοιχεία των θέσεων Υ και Χ, αυτός ο τουρμαλίνης μπορεί να θεωρηθεί ως μετα-μαγματικής γένεσης, πιθανά με τη συνδρομή υδροθερμικών ρευστών.

Τουρμαλίνης, κρυσταλλοχημεία, ακραία μέλη, ουβίτης - φερουβίτης, ποβονδραΐτης, πηγματίτης, Σιθωνία, Ελλάδα.

1. Introduction

Tourmaline is a complex cyclosilicate belonging to the hexagonal crystal system, with general formula XY3Z6[BO3]3T6O18V3W. Its variable chemism, frequent zoning, wide stability field P-T, ability to exchange chemical species with coexisting minerals and fluids (Keller et al. 1999) and high strength to exogenous agents make this mineral a potential metallogenetic and petrogenetic indicator in many igneous and metamorphic environments (Manning 1982, Henry and Guidotti 1985, Jolliff et al. 1986, Kassoli-Fournaraki 1990, Farmer and Halls 1993, London and Manning 1995, Henry and Dutrow 1992, 1996).

Tourmaline structure consists of five types of cationic site: T, tetrahedrally coordinated; B, in triangular coordination; Ζ and Y, octahedrally coordinated; X, coordinated by nine or ten anions (Selway 1999). Both two octahedral sites and X site are involved in various possible substitutions which give rise to the wide compositional range showed by natural tourmaline.

Usually Al is the dominant cation in Ζ site, sometimes substituted by even considerable amounts of Mg, Fe2+, Fe3+ (Fe3+6 = oxy-ferri-foitite and oxy-Mg-ferri-foitite), Cr (Cr6 = chromdravite and fluor-chromdravite); since substitutions at Y site are constrained more by the size of the ions than by their valence (Grice and Ercit 1993), many different cations may dominate it, that is Li (Li2Al = liddicoatite and hydroxy-liddicoatite), Mg, Fe2+, Fe3+, Al and Cr (MgCr2 = oxy-chromdravite) (Hawthorne and Henry 1999), with lesser amounts of Mn, V, Zn (Schmetzer and Bank 1984, Kazachenko et al. 1993).

Several cations may get into X site, mainly occupied by Na, Ca; moreover, it can be K-dominant (Grice et al. 1993) or show remarkable vacancies (MacDonald et al. 1993, Pezzotta et al. 1996, Selway et al. 1999a, Aurisicchio et al. 1999).

The sites occupied by Si and B are subject to moderate chemical substitutions: small amounts of Al substitute for Si cations in TO4 tetrahedra (Foit and Rosenberg 1977, 1979, Hawthorne 1996). The presence of B inside TO4 is debatable, some authors assert there is no evidence for its occupation of Si tetrahedral site, whereas recently Erl et al. (2006) describe tourmaline...
containing IVB from several localities. The presence of substantial IVB is limited to, or more common in, Al-rich tourmalines.

The anions occurring in V and W sites allow defining if the observed phase is an oxy, a hydroxy or a fluoritic term: V site is usually filled by OH\(^-\), rarely by O\(^2-\); W site may be occupied by O\(^2-\), OH\(^-\) or F\(^-\).

Two main mechanisms of substitution can occur in tourmaline: the first one involves changes of cations having the same oxidation state inside the same crystallographic site (i.e. \(\text{Y}^\text{Mg} \leftrightarrow \text{Y}^\text{Fe}\text{Fe}^{\text{3+}}\)); the second one involves changes of heterovalent cations between different sites (i.e. \(\text{X}^\text{Ca} + \text{Z}^\text{Mg} \leftrightarrow \text{X}^\text{Na} + \text{Z}^\text{Al}\)) (Henry and Guidotti 1985).

Tourmaline composition was currently expressed on the basis of its end-members dravite, schorl and elbaite. However, the analytical method evolution and the new impulse given to researches on this mineral due to its genetic significance showed remarkable deviations and the opportunity to consider new end-members for every possible succession in the three tourmaline groups: alkali, calcic and X-vacant (Hawthorne 1999).

In this work, we discuss the features and the genetic inferences of small tourmaline grains occurring in pegmatitic dykes cutting metamorphic rocks of the Circum-Rhodope Belt, an isopic zone of the internal Hellenides, and related to the nearby granitic intrusions of Sithonia peninsula. The tourmalines studied occur within the dykes as swarms of clusters, composed by small grains or mostly tiny subhedral crystals, showing under microscope evidences of cracks, rarely slight traces of bending, homogeneous distribution of the color (parallel polars) and a sharp pleochroism from pale yellow to deep green. Both the granular aggregates and the subhedral crystals show similar composition.

2. Geological Environment

The Sithonia Peninsula, the middle extension of Peninsula (Northern Greece), is occupied by rocks belonging to the Circum-Rhodope Belt (CRB) and the Serbo-Macedonian Massif (SMM). The CRB is mostly covered by the Sithonia pluton, an elongated Eocene intrusion, and in lesser extent by a Permian to Jurassic flysch of the Svoula Group, by Mid-Jurassic Ophiolites, and by the Chortiatis Unit (Fig. 1) consisting of greenschists, gneisses, metadiorites and marbles (Kauffman et al. 1976, Kockel et al. 1977). The SMM rocks of the area include the Palaeozoic metamorphic rocks of the Vertiskos Series (Fig. 1) (Kockel et al. 1977).

The Sithonia pluton displays irregular shape and follows the NW-SE regional structural trend. In many places, particularly at the margins of the pluton, a clearly oriented fabric and cataclastic texture can be observed, very probably related to a regional deformative and metamorphic event. The intrusive body comprises moderately peraluminous to metaluminous rocks; the geochemical features suggest processes of fractional crystallization and crustal contamination for the genesis of the Sithonia pluton (Christophides et al. 1990).
The ophiolitic rocks and the Chortiatis Unit rocks have been deformed and regionally metamorphosed in the upper greenschist to amphibolite facies mainly during Upper Jurassic to Lower Cretaceous. Retrograde metamorphic events have been also recognized (Mussallam and Jung 1986, Kassoli-Fournaraki 1990).

Aplitic and pegmatitic veins cut both intrusive and country metamorphic rocks.

The samples taken for this study are representative of a dyke cross-cutting gneissic rocks of the Chortiatis Unit in the central area of the peninsula (Fig. 1).

3. Experimental

Chemical compositions were determined in the Sezione di Roma of the Institute of Geosciences and Georesources of the CNR, using a Cameca CX 827 Electron Microprobe, equipped with three wavelength-dispersive (WDS) and one energy-dispersive (EDS) Link-System spectrometers, which are running simultaneously.

Selected crystals have been analyzed for a definite group of elements distinctive of tourmaline, among which only Si, Al, Na, Ca, Mg, Fe, Ti and Mn exceed instrumental detection limits. Compositional homogeneity was verified within single crystals and between crystals belonging to different swarms, to specify eventual chemical variations.

Operating conditions were 15 KV accelerating voltage, 20 nA beam current and 2μ beam diameter. Wavelength spectrometers were used for all analyzed elements; element peaks and background were measured with counting times of 10 seconds each. A synthetic diopside glass was used as a reference material for Si (TAP), Ca (PET) and Mg (TAP); a jadeite crystal for Al (TAP) and Na (TAP); a rutile crystal for Ti (PET); a Norway fayalite for Fe (LIF); rhodonite for Mn (LIF); a Durango apatite for F (TAP); orthoclase for K (PET). Matrix correction was calculated by using the \( \varphi (pZ) \) method (Pouchou and Pichoir 1984, 1985).

EMP analyses representative of compositions of both granular aggregates and subhedral crystals are reported in Table 1, together with the number of atoms per formula unity (apfu) calculated on the basis of 31 (O, OH), assuming stoichiometric amounts of OH\(^-\) and B. The assumption B=3 apfu is sustained by several authors (Grice et al. 1993, Hawthorne et al. 1993, Selway et al. 1998) and seems to be satisfactorily correct. Assessment of water content also could be well represented by the formula OH=4apfu (MacDonald et al. 1993, Selway et al. 1998). The Li content was obtained on the basis of the formula: 3-ΣY. Fluorine was always under detection limits.

Crystallographic sites have been filled partially following previous works (Grice and Ercit 1993, Frietsch et al. 1997) and as indicated by some clear correlations.

In order to estimate the \( \text{Fe}^{2+}_{\text{tot}}/\text{Fe}^{3+}_{\text{tot}} \) ratio, we imposed the saturation of Si tetrahedral site (Si=6.000 apfu). Using Grice and Ercit (1993) formula the number of Mg atoms at Y site has been calculated; Z site has been saturated to 6.000 apfu with any remaining Mg ions and Fe considered as Fe\(^{3+}\) (Frietsch et al. 1997).

4. Mineral Chemistry Data and Discussion

The Sithonia tourmaline is characterized by Al deficiency (Al<6.0 apfu), abundance of Mg coupled with high Fe and prevalence of Na in X site; it has an unusual composition, seldom occurring in tourmaline literature, ascribed by some authors to mere dravitic members (Grice and Ercit 1993) or as povondraitic oxy-dravite (Henry et al. 1999).

Considering Table 1, Al content is never sufficient to fill the octahedral Z site, whereas Y site occupancy is greater than 3 apfu as was expected considering that some Mg and/or Fe should enter Z site to fill it. Mg ranges from 1.749 to 2.032 apfu, always exceeding \( \text{Fe}^{2+}_{\text{tot}} \) which ranges from...
1.050 to 1.597 apfu. Mn is usually under the instrumental detection limits, while Ti sometimes reaches 0.14 apfu. Calculated Li is never greater than 0.15 apfu, as expected for ferro-magnesian compositions.

Tetrahedral T-site always shows a considerable excess of Si cations (Table 1); this repeated result has been confirmed by rigorous checks through the execution of numerous analyses over different crystals and within single crystals and using different reference standards, in order to display Q-micro-inclusions effect and eventual instrumental errors. This excess, as consequence, could be considered consistent with the existence of trivalent cations in the tourmaline structure.
Figure 2 - Linear correlations between the main elements (apfu) of Sithonia tourmaline: (a) Fe\(^{3+}\) tot vs. Mg\(_{\text{tot}}\); (b) Na vs. Al\(_{\text{tot}}\); (c) Ca vs. Al\(_{\text{tot}}\); (d) Mg\(_{\text{tot}}\) vs. Al\(_{\text{tot}}\)

X-site is dominated by Na, followed by lesser amounts of Ca. Small vacancies are frequent.

The following linear correlations between the main elements allow to visualize the Sithonia tourmaline behavior: Fe\(^{3+}\) tot vs. Mg\(_{\text{tot}}\) shows oscillations around a trend slightly negative (Fig. 2a); a light positive correlation links Na to Al\(_{\text{tot}}\) (Fig. 2b), whereas Ca and Al\(_{\text{tot}}\) give a good negative correlation (Fig. 2c); moreover, Mg shows steady values for all Al\(_{\text{tot}}\) contents (Fig. 2d).

Figure 3 - (a) Plot of Ca/(Ca+Na) at the X site vs. Fe/(Fe+Mg) at the Y site. (b) Al – Fe\(_{\text{tot}}\) – Mg diagram (in molecular proportions) for tourmaline from various rock types: field #6 defines the compositional range of tourmaline from Fe\(^{3+}\)-rich quartz-tourmaline rocks, calc-silicate rocks and metapelites (Henry and Guidotti 1985)

The diagrams of Fig. 3 (a and b) represent an attempt for the classification of Sithonia tourmaline on the basis of the correlation Ca/(Ca+Na) vs. Fe\(_{\text{tot}}\)/(Fe\(_{\text{tot}}\)+Mg) given by Selway et al. (1998) and the compositional fields defined in the triangular plot Al\(_2\)O\(_3\) – MgO – FeO proposed by Henry and Guidotti (1985). In Fig. 3a the studied compositions fall on the boundary line between schorl and dravite end-members, with a light prevalence of dravite. The triangular diagram of Fig. 3b shows the area covered by Sithonia compositions: they fall below the Al\(_{50}\) – (Fe, Mg)\(_{50}\) line.
However, the reported graphic representations of the chemical compositions of Sithonia tourmaline do not seem to be fully representative of these phases, since they do not show meaningful trends and do not allow to point out the Al behavior, nor to define any specific field for such compositions. Nevertheless, the triangular diagram of Fig. 3b suggests the presence of a possible considerable Fe$^{3+}$ amount, since they fall just below the field #6 of the triangle reported by Henry and Guidotti (1985) and assigned to Fe$^{3+}$-rich quartz-tourmaline rocks, to calcium-silicate rocks and to metapelite. This hypothesis may be further confirmed by the optical observations carried out on our samples. In fact, tourmaline containing significant amounts of Fe with mixed valences (2+ and 3+) exhibits very strong optical absorption and pleochroism due to intervalence charge transfer (Mattson and Rossman 1987), as in the case of Sithonia tourmaline.

Figure 4 - Tendency of calculated Fe$^{3+}$ vs. Al$_{tot}$ showing a negative correlation (apfu)

To quantify the amount of Fe$^{3+}$ the Fe$^{2+}_{tot}$/Fe$^{3+}_{tot}$ ratio was evaluated, using an iterative program, which carries out Fe$_{tot}$ division imposing a stoichiometric amount of Si=6,000 apfu. In Table 2 the analyses of a single tourmaline crystal have been recalculated following the above procedure.

Using the formula proposed by Grice and Ercit (1993), based on the hypothesis of an ideal solid solution at Y site between Fe and Mg for tourmaline compositions with FeO$_{tot}$ greater than 7 %, it was possible to determine the number of Mg atoms at Y site, with any remaining Mg assigned to Z. Often the calculated $^5$Mg is not enough to fill Z site and vacancies continues to exist which have been saturated with the necessary amount of Fe$^{3+}$ to reach 6,000 apfu.

The goodness of this choice, to partially fill Z-site with Fe$^{3+}$, is confirmed by the good negative trend Fe$^{3+}$ vs. Al$_{tot}$ (Frietsch et al. 1997) of Fig. 4, the excellent negative trend of Fig. 5 shows the uvitic substitution Na$^+$/Al$^3+$ $\leftrightarrow$ Ca$^+$/Mg. Moreover, the Manning (1982) diagram points out transitional compositions toward uvite end-member perfectly following substitution mechanism line (Fig. 6). So, the seemingly simple solid solution schorl-dravite reveal the existence also of an

Figure 5 - Negative correlation Na+Al(Z) vs. Ca+Mg(Z), representing the uvitic substitution (apfu)

Figure 6 - The complex substitution mechanism Fe$^{3+}$+Na+Li $\leftrightarrow$ Mg(Y)+Ca+Fe$^{2+}$, controlling the transition povondraite $\rightarrow$ uvite-feruvite (apfu)
uvitic component, as from the entrance of Ca substituting Na in X site proceeding toward lesser Al contents, with always high Mg$_{tot}$.

Assuming the Al amount as indicator of compositional and thermal evolution, it is known that Al-richer tourmaline crystallize during magmatic and pegmatitic stages, while Al-poor terms have been observed in the later hydrothermal stage (London and Manning 1995, Trumbull and Chaussidon 1999), Sithonia tourmaline crystallized in a Fe-Mg rich and Al-poor environment, according with the fluorine content always under the electron microprobe detection limits, which, coupled with the decrease of T, should expand the stability field of tourmaline toward lower B and Al content (London 1999).

The recorded Fe$^{3+}$ contents are variable but increasing for decreasing Al (Table 2), therefore while tourmaline growth evolves from schorl-dravite toward the uvitic end-member, the oscillatory trend of Fe$^{3+}$-rich component is the result of a fluctuating increase of the oxydation degree of hydrothermal fluids. Because of the increase of oxygen fugacity, the transition from simple schorl-dravite to more complex solid solution could be combined also with the substitution of OH$^-$ with O$_2^-$ in W site. A theoretical evaluation of Bond Valence Sum (BVS), utilizing ionic radii values of Shannon (1976) and Brown and Wu formula (1976), with the updating of Brown (1987), allows confirming the predominant OH$^-$ presence in W site, consistently with the results obtained by Grice and Erce (1993) for similar tourmaline compositions on the basis of X-ray data.

In order to emphasize the significant chemical features of these phases, we consider a ternary diagram functional to support the new reading-key, that is Al$_6$ – Fe$^{3+}$$_6$ – Mg$_6$ (Fig. 7a), which three vertices describe Z-site occupancy: the first Al$_6$ indicates the Al saturation amount both in dravite and scorial end-members; the second Fe$^{3+}$$_6$ the Fe$^{3+}$ saturation amount in an oxy-Mg-ferrofoitite and an oxy-ferroferrite, while Mg$_6$ is the Mg amount in a hypothetical term in which Mg entirely fills Z-site. For the sake of clearness, in the same diagram some Z-compositions are plotted, representative of tourmaline end-members particularly useful to display Sithonia tourmaline behavior: the representative points of Sithonia tourmaline fall near the line jointing the top vertex Al$_6$ with MgAl$_5$ point corresponding with Z occupancy of (hydroxy)uvite, (hydroxy)feruvite and oxy-dravite. The diagram shows that the Al deficiency causes a shift towards “uvite” composition, because the existence of an oxy-dravitic component can be excluded owing to Al(Y) absence. Moreover, the F absence in the analyses leaves out Mg Al$_5$ point represents compositions with F anions in W-site (i.e. ideal uvite and feruvite).

![Figure 7](http://epublishing.ekt.gr)

**Figure 7** - (a) Triangular diagram Al$_6$ – Fe$^{3+}$$_6$ – Mg$_6$, describing Z-site occupancy, (b) Triangular diagram Fe$^{3+}$$_3$ – Mg(Y)$_3$ – Fe$^{2+}$$_3$, showing the coexistence of the three components
In the light of the above, we can consider Sithonia compositions as schorl-dravite containing an uvitic and a Fe\(^{3+}\)-rich component. The triangular diagram of Fig. 7b, which vertices represent the saturation amounts of Fe\(^{3+}\) (Fe\(^{3+}\)-rich end-member), Mg (dravite) and Fe\(^{2+}\) (schorl) in Y-site, defines the compositional field of this solid solution.

The small dimensions of tourmaline crystals, together with their exclusive occurrence as subparallel veinlets, as well as their complex composition characterized by Fe\(_{tot}/Mg\(_{tot}\) ratio<1 (Power 1968, Fieremans and De Paepe 1982, Taylor and Slack 1984, London and Manning 1995, Trumbull and Chaussidon 1999), suggest that tourmaline crystallized during a hydrothermal stage, in subsolidus conditions, as confirmed by \(^{27}\)Al absence (London and Manning 1995) and by Al deficiency in Z site. The low degree of alkali-site vacancies (Trumbull and Chaussidon 1999) allows drawing similar inferences on the crystallization of Sithonia tourmaline.

5. Conclusions

- Compositions as Sithonia tourmaline, before considered as simple dravite-schorlrite solid solutions or povondraitic oxy-dravite, turn out to be a more complex phase, characterized by deficiency of Al in Z-site and predominance of Mg\(_{tot}\) over Fe\(_{tot}\), with Na over Ca in X-site. This composition seems to refer to an uncommon combination of several actual and hypothetical end-members.

- The hypothesis of a hydrothermal origin is further on supported by the finely oscillatory trends, distinctive of a growth from rapidly changing fluids. The several peaks of Fe\(^{3+}\) within its negative oscillatory trend versus Al\(_{tot}\) could be referred to different and more and more oxidized fluid rising (Cavarretta and Puxeddu 1990).

- The variations of the chemical-physical parameters produced a fine-scale zonation of crystals, reflected by continual trend inversions (London 1999). At this state, lacking isotopic studies, it is not possible any claim about the origin of fluids: true metamorphic or with a magmatic component. However, a strong metamorphic contribution seems to be preferred on the basis of micro-structural and geological observations, i.e. the bending traces under the microscope and the gneissic rocks surrounding the quartz-tourmaline dyke; moreover, we cannot avoid to remark that hydrothermal tourmaline, containing high Mg, usually occurs in veins hosted in metasedimentary rocks (Smith and Yardley 1995).

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7. References


Manning, D.A.C., 1982. Chemical and morphological variation in tourmalines from the Hub Kapong batholith of peninsular Thailand, Mineralogical Magazine, 45, 139-147.


