

GEOLOGICAL, GEOMORPHOLOGICAL AND TECTONIC STRUCTURE OF NE ATTICA AND SEISMIC HAZARD IMPLICATIONS FOR THE NORTHERN EDGE OF THE ATHENS PLAIN

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Abstract

A synthesis of geology, geomorphology and tectonics has been compiled regarding the NE part of Attica. This synthesis helps us clarify how old and new structures interrelate and interact to provide the present day setting. Geological, geomorphological maps, and cross-sections are provided to help us depict and extract data. The region of NE Attica forms a tilted tectonic block bounded by the Afidnai fault to the south and the Oropos fault to the north that rotates to the S-SW. This tilt produces southern trending flow directions draining the footwall within the block. Drainage basins are highly asymmetric due to the presence of active normal faults producing a combination of fault parallel and fault perpendicular flow directions. This block is also divided by a NNE-SSW detachment fault that separates the metamorphic units to the east from the unmetamorphic units to the west. It was active in Late Miocene-Early Pliocene and produced several hundred meters of debris-flow deposits. This detachment influences the geometry, style and intensity of deformation, but also the seismicity pattern. In particular, this detachment coincides with the line separating zone I (lowest category of seismic risk) from zone II of the national seismic building code. Finally, the Athens plain is bounded northwards by the active, but low slip-rate E-W trending, 14 km long, Afidnai fault.

Key words: Detachment, active fault, Afidnai, Oropos, Marathon.

Περίληψη

Με βάση τη σύνθεση των γεωλογικών, γεωμορφολογικών και τεκτονικών δεδομένων της ΒΑ Αττικής γίνεται διάκριση της αλληλοσυσχέτισης μεταξύ παλαιών και νέων δομών που διαμορφώνουν την παρούσα κατάσταση. Η εξαγωγή δεδομένων βασίστηκε σε γεωλογικούς και γεωμορφολογικούς χάρτες και τομές. Η περιοχή της ΒΑ Αττικής αποτελεί ένα ρηξιτέμαχος το οποίο οριοθετείται από το ρήγμα των Αφιδνών στα νότια και το ρήγμα του Ωρωπού στα βόρεια και περιστεύεται προς τα νότιο-νοτιοδυτικά. Η περιστροφή αποστραγγίζει το υδρογραφικό δίκτυο προς το νότο πάνω στο δάπεδο του ρηξιτέμαχος. Οι λεκάνες του υδρογραφικού δικτύου είναι έντονα ασύμμετρες λόγω

της παρουσίας των ενεργών ρηγμάτων. Οι διευθύνσεις ροής των ρεμάτων λόγω της δράσης των ρηγμάτων είναι είτε παράλληλες είτε εγκάρσιες ως προς τα ρήγματα. Το ρηξιτέμαχος διαιρείται μέσω ενός BBA-NNΔ ρήγματος αποκόλλησης σε ένα ανατολικό τμήμα με μεταμορφωμένα πετρώματα και σε ένα δυτικό με αμεταμόρφωτα. Το ρήγμα αυτό ήταν ενεργό κατά το Ανώτερο Μειόκαινο – Κατώτερο Πλειόκαινο και μέσω της δράσης του προκάλεσε την απόθεση πολλών εκατοντάδων μέτρων αδιαβάθμητων ιζημάτων κροκαλο-λατοποπαγών. Το ρήγμα αυτό επηρεάζει τόσο την γεωμετρία, το είδος και την ένταση της παραμόρφωσης όσο και την σεισμικότητα της ευρύτερης περιοχής. Πιο συγκεκριμένα, το ρήγμα αποκόλλησης ταυτίζεται με το όριο που διαχωρίζει την κατώτερη κατηγορία σεισμικότητας I, από την ενδιάμεση κατηγορία σεισμικότητας II του Ελληνικού Αντισεισμικού Κανονισμού (ΕΑΚ-2003). Τέλος, το Λεκανοπέδιο Αθηνών οριοθετείται προς βορρά από το ενεργό, αλλά χαμηλού ρυθμού ολίσθησης, ρήγμα των Αφιδνών μήκους 14 χλμ.

Λέξεις κλειδιά: Ρήγμα αποκόλλησης, Ενεργό ρήγμα, Αφίδνες, Ωρωπός, Μαραθώνας.

1. Introduction

The northeastern part of Attica forms a complex mosaic of old and new structures. These structures shape the present-day geomorphological, neotectonic and seismicity status. In this paper we: a) try to unravel the geological history of the area since the Late Miocene, b) show how these structures interrelate and interact and c) discuss the implications for seismic hazard assessment.

2. Geological Structure

2.1. Geotectonic Units, Alpine rocks and post-Alpine sediments

Northeastern Attica comprises Alpine basement rocks, both metamorphic and non-metamorphic, and post-Alpine sediments (Fig. 1). The Alpine rocks belong to the metamorphic units of the Northern Cyclades and Almyropotamos that extend from Penteli Mt to the southern Evoikos gulf and to the non-metamorphic unit of Eastern Greece that outcrop in Parnitha Mt (Papanikolaou 1986). The Almyropotamos unit is known from southern Evia, but it crops up also in the area of Marathon. It comprises a thick Mesozoic marble sequence overlain by phyllites representing a metamorphosed Tertiary flysch (Katsikatsos 1969, Dubois and Bignot 1979). The Northern Cyclades unit emerges not only on the northern Cyclades islands, but also in a major part of the southern Evia, where it is known as the Styra unit (Katsikatsos 1979). It consists of thick sequences of mica schists and cipolines grading to siliceous marbles with alternations of greenschists and amphibolites, representing a pelagic volcanosedimentary sequence. More to the south the Northern Cyclades unit is tectonically overlying the relative autochthon unit of Attica (Lozios 1993). The non-metamorphic tectonic unit of Eastern Greece comprises: (i) the SubPelagonian Triassic-Jurassic carbonate platform overlying a volcanosedimentary Permo-Triassic complex and underlying an upper Jurassic schist-hornstein melange formation, (ii) the ophiolite nappe of the Vardar/Axios oceanic basin, and (iii) the upper Cretaceous transgressive sequence made of a shallow-water carbonate platform and a Tertiary flysch. Most part of the Parnitha Mountain consists of the lower group (i). The deformation of the Alpine formations of the above units is post-Late Eocene, as dated from the age of the flysch involved in the tectonic structures and extends up to the Early Miocene as the synmetamorphic deformation shows in the metamorphic rocks of the Cyclades (e.g. Schliestedt *et al.* 1987).

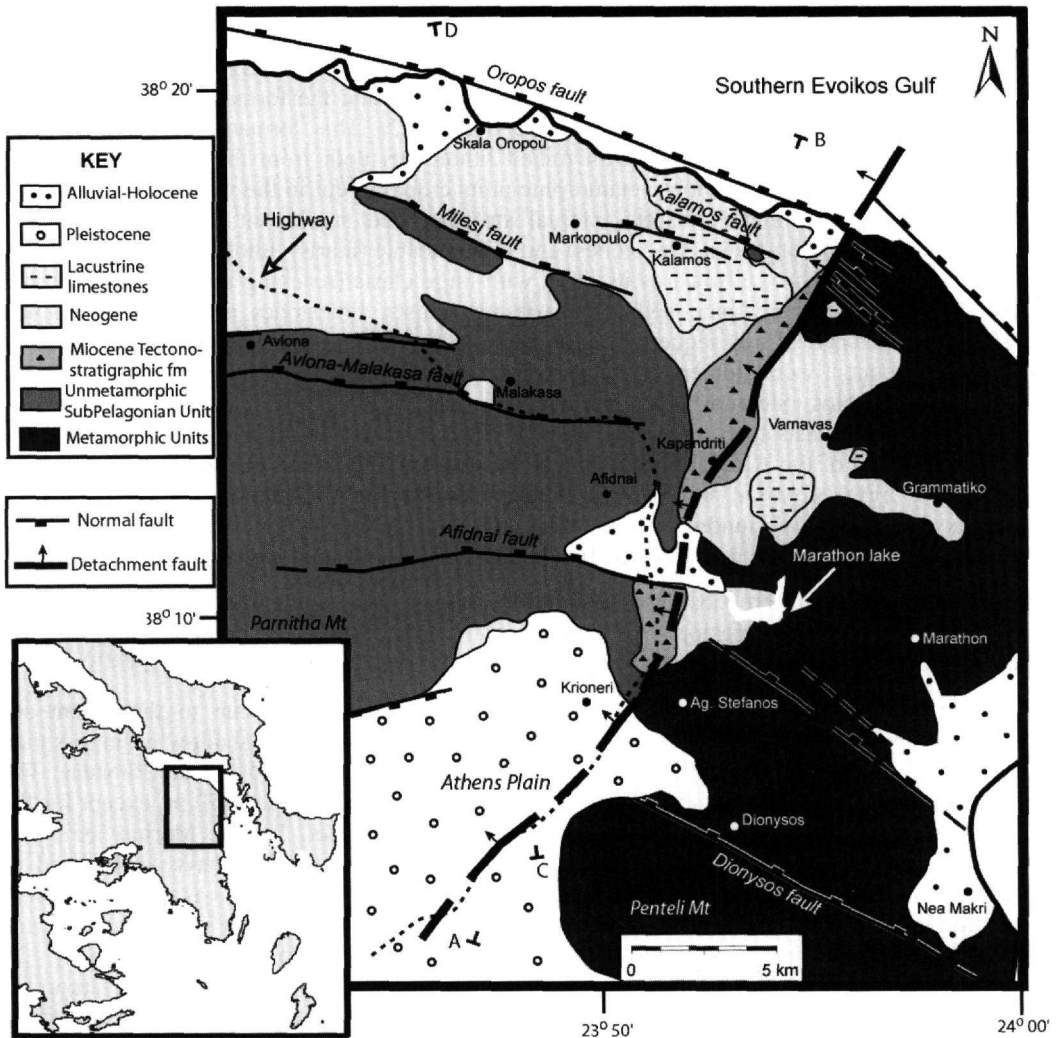


Figure 1 - Geological map of NE Attica

A detachment fault (see below 2.2) separates the metamorphic units from the non-metamorphic units. This detachment passes from the Ochtonia cape in the Aegean coast of Southern Evia (northwards our study area), through Aliveri to Kalamos in northeast Attica and continues to the southwest into the plain of Athens approximately along the Kifissos River (Papanikolaou *et al.* 1999, Xypolias *et al.* 2003). Finally, between the metamorphic and the non-metamorphic units and along strike the tectonic contact there is a narrow zone, where some small outcrops of low-grade metamorphic rocks occur under the post-Alpine sediments that cannot be correlated to either of these two units. These rocks might be correlated to the Alepovouni unit or a similar type of unit observed further south (Papanikolaou *et al.* 2004a), whose lower part comprises a few hundred meters of semi-metamorphic sandstones, schists, and phyllites. The Alepovouni unit is bounded by two detachment faults one separating it from the overlaying unmetamorphic units of Athens and the SubPelagonian and the other detachment from the underlying metamorphic Attica autochthon unit (Papanikolaou *et al.* 2004a).

The post-Alpine sediments of the area comprise Late Miocene – Pliocene continental deposits and minor outcrops of Pleistocene and Holocene alluvial (e.g. Mettos 1992). The older sediments comprise a clastic tectonosedimentary formation that outcrops in the area of Kapandriti, and in the area between Aghios Stefanos and Afidnai (Figs 1, 2, 4). They form a NE-SW zone of debris-flow deposits that are several hundred meters thick (Fig. 3), indicating the formation of a major fault zone in a steep relief during Late Miocene (see below 2.2). This formation consists of large boulders and blocks in a clay matrix with mixed lithologies both from the metamorphic units cropping out in the east and the non-metamorphic units cropping out in the west. In Kapandriti, it is characterized by a complex/chaotic internal structure and significant lateral variations of permeable (sands, conglomerates, sandstones) and impermeable (marls) lithologies. Semi-cohesive breccio-conglomerates prevail towards its base, with large clasts (reaching up to a diameter of 2 meters) sourced from the alpine bedrock involving schists, limestones and cherts (Fig. 4a). Towards the upper part of this sequence, clasts become thinner and finer and at places sandstone beds and less frequently conglomerates are also observed. The same formation outcrops also in the area between Aghios Stefanos and Afidnai and has been evolved laterally to breccio-conglomerates with clasts that are smaller in size (rarely exceed 30-40 cm in diameter). In the immediate footwall of the Afidnai fault its thickness exceeds 150 m and the bedding dips 20°-40° towards NNE (Fig. 4b). On top of this breccio-conglomerate formation outcrop lacustrine limestones that are a few hundred meters thick. The top lacustrine limestones are also found unconformably overlying the metamorphic basement in the east around Varnava.

Pleistocene alluvial fans are observed within the Athens plain along the southern slopes of Parnitha Mt and the western slopes of Penteli southwards from the Afidnai and the Aghios Stefanos, unconformably covering the Neogene sediments (e.g. Papanikolaou *et al.* 2004b).

No outcrops of marine sediments are traced in Northeastern Attica, due to the recent sea transgression in the southern Evoikos gulf. The southern Evoikos gulf is a shallow basin less than 250 m deep that separates Attica from the southern Evia and was formed in Late Pliocene. The thickness of the Plio-Quaternary sediments within the gulf does not exceed 150 m, except for the southeastern area, where they are 250 m thick (Papanikolaou *et al.* 1988a). Marine sediments occur only along the coast of eastern Attica at Rafina area where they have been dated as uppermost Pliocene – lower Pleistocene (Mitzopoulos 1948, Guernet and Sauvage 1970). Finally, recent alluvial sediments are limited in thickness and are observed only in some small narrow bands along the coastline (Oropos, Kalamos, Marathon) and in the Afidnai basin (Fig. 2). Roubanis (1961) reported that in the Afidnai plain close to the railway station, the metamorphic basement was drilled at a depth of 47 m and that the Neogene sediments were only 15 m thick, revealing that the Late Pleistocene-Holocene sediments are about 30 m thick. Recent drillings for water in the Plain also showed that the alluvial- Quaternary thickness does not exceed 80 m.

2.2. The role of the detachment fault

The dominant structure is a major NNE-SSW trending detachment, separating the metamorphic units towards the east from the non-metamorphic units towards the west. Moreover, it separates the E-W trending faults in the western part from the NW-SE faults in the eastern part. This detachment caused the downward movement of the non-metamorphic units and the uplifting of the metamorphic units from the deeper part of the lithosphere where the metamorphism took place in Eocene-Oligocene times. The Miocene tectonosedimentary sequence of debris-flow deposits rests on the hangingwall of the detachment fault and was formed by its activity. This extensional detachment was active throughout Late Miocene times and gradually became inactive during Early Pliocene, when lacustrine deposits have been accumulated all over the area from Varnava and Kapandriti to Kalamos (Fig. 1). The same detachment can be traced further south. In particular, it has subsided the non-metamorphic SubPelagonian unit towards the NNW and uplifted the Alepovouni and the Attica autochthon to the SSE (Papanikolaou *et al.* 2004a).

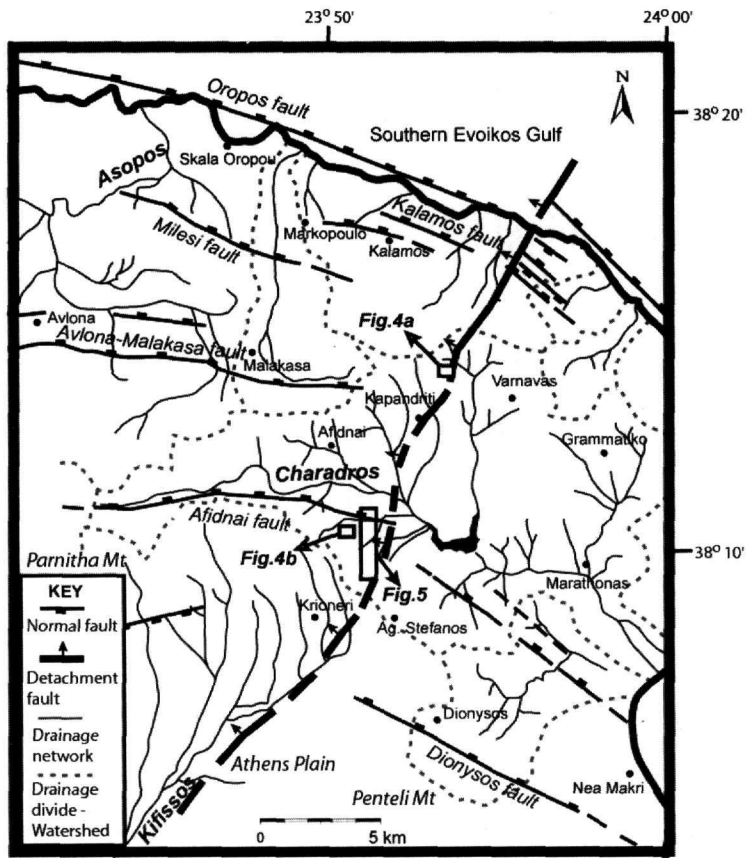


Figure 2 - Drainage network, drainage divide and active faults in NE Attica

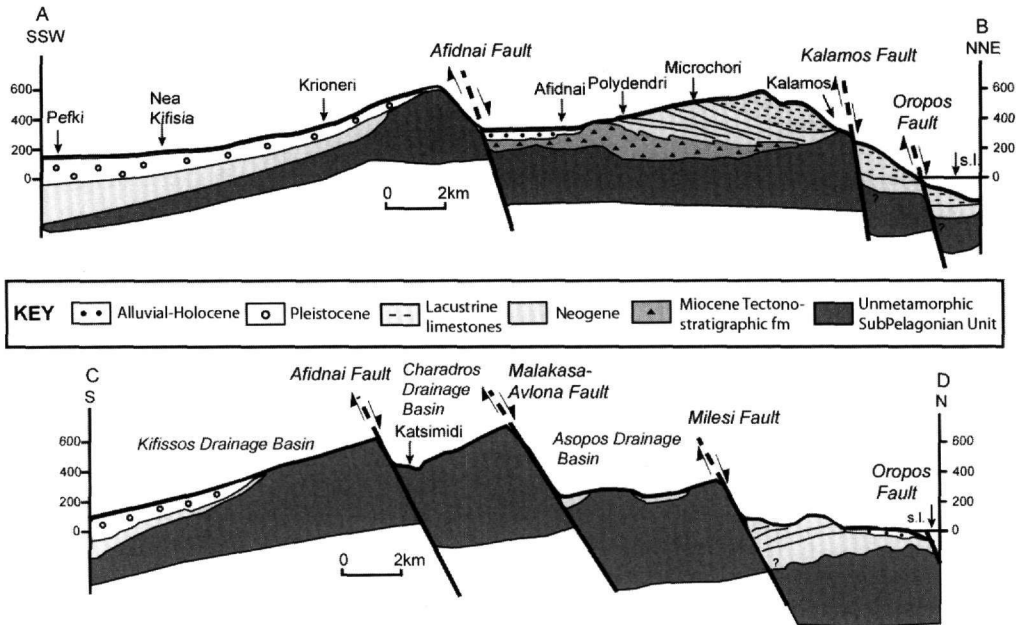


Figure 3 - Cross-sections A-B and C-D (see Fig. 1)

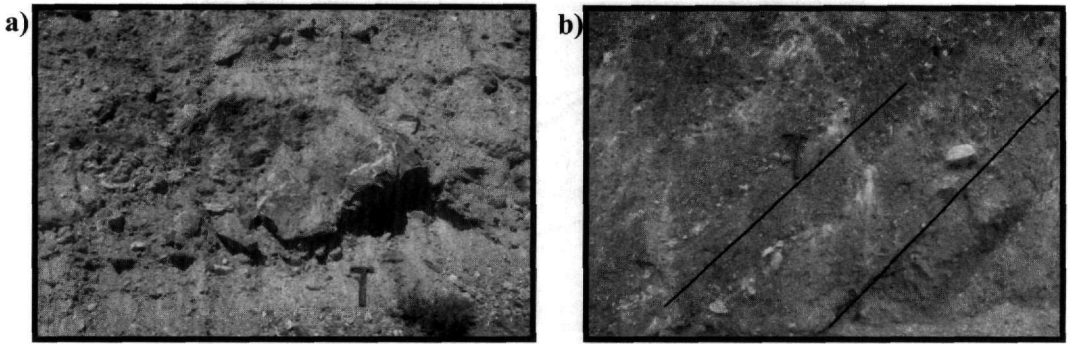


Figure 4 - a) View of the Miocene tectonosedimentary formation of debris-flow deposits in the area of Kapandriti (for location see Fig. 2). Clasts are large (up to a diameter of 2 m), implying that this locality was very close to the detachment trace. b) View of the same formation in the area between Aghios Stefanos and Afidnai in the immediate footwall of the Afidnai fault. The bedding plane is nicely revealed, dipping 30°-40° towards NNE. Clasts are relatively small in size rarely exceeding 30-40 cm in diameter, indicating that this locality was less proximal to the detachment

3. Geomorphological and neotectonic structure

The geomorphology of the area is dominated by the Penteli and Parnitha mountains, which border the Athens Plain from the north and the east, respectively. Further north from the southern Evoikos gulf and the coastal zone of Oropos-Kalamos up to the Athens Plain, three major drainage basins are observed: 1) The Kifissos basin in the south with a NNE-SSW flow direction, draining the Athens plain from the southern slopes of Parnitha and the southwestern slopes of Penteli to the Saronic gulf in the southwest. This flow direction is parallel to and near the detachment trace; 2) The Charadros basin with a major W-E flow direction, sourced from the eastern Parnitha highs of about 1200 m up to the lows of the Afidnai plain and the Marathon lake at about 350 m; 3) The Asopos basin towards the northwest which flows northwards, but has a more complex structure due to the existence of the Avlona-Malakasa and the Milesi faults. Finally, some small approximately N-S trending catchments are observed in the area of Kalamos and the Microchori plateau from 500 m elevation up to the coast of southern Evoikos.

Drainage basins are highly asymmetric due to the presence of active normal faults. In particular, there is a combination of fault parallel and fault perpendicular flow that is characteristic of active normal faulting settings (Gawthorpe and Hurst 1993, Eliet and Gawthorpe 1995). The tilted footwall results into fault perpendicular flow directions, which drain the footwall away from the hangingwall (e.g. footwall of the Afidnai fault, footwall of the Oropos-Kalamos fault in the area of Kapandriti; see also cross section Fig. 3). On the other hand, headward erosion within the footwall catchments drain across the fault into the hangingwall producing also fault perpendicular flow directions (e.g. Avlona-Malakasa fault, Afidnai fault, Oropos fault). Finally, several catchments are clearly deflected into a fault parallel flow direction due to hangingwall subsidence (e.g. Afidnai fault, Avlona-Malakasa fault, Milesi fault).

The planation surfaces, both erosional and sedimentary, from 50 m up to 1000 m in altitude, show an E-W preferred orientation and are usually inclined southwards, indicating a tilt around an E-W axis (Papanikolaou *et al.* 1988b). This tilt is nicely illustrated in cross-sections (Fig. 3).

Active faults are trending E-W in the northwestern part and NW-SE in the southeastern part; a common tectonic pattern along the Hellenic arc (Mariolagos and Papanikolaou 1981, 1987). The E-W trending faults are large and known to generate medium to large magnitude earthquakes (1938 Oropos M=6.0, 1981 Alkyonides earthquake sequence M=6.7, M=6.4, M=6.3, 1999 Athens M=5.9), whereas the NW-SE faults occurring in the east are shorter, have lower rates and produce

low to medium magnitude earthquakes (Papanikolaou and Lozios 1990). This neotectonic pattern has been analysed in the northern Attica through morphotectonic observations (Papanikolaou *et al.* 1988b), as well as offshore studies in the Saronikos and the southern Evoikos gulfs (Papanikolaou *et al.* 1988a).

This set of E-W trending north dipping active normal faults in NE Attica consists of: 1) the Oropos-Kalamos fault that is parallel to the coast of the southern Evoikos gulf. It is at least 14 km long and part of it probably ruptured during the M=6.0 event in 1938. The southern Evoikos gulf is a shallow basin less than 250 m deep and the main fault of the Gulf has a throw of 250 m (Papanikolaou *et al.* 1988a), implying that the Oropos-Kalamos zone as a whole is a relatively small structure and of similar importance as the other across strike faults; 2) the ~10 km long Milesi fault, which produces a characteristic bedrock scarp on the Mesozoic limestones that form the footwall and the continental Neogene sediments resting on the hangingwall; 3) the ~ 18 km long Avlona-Malakasa fault that borders the northern slopes of Parnitha Mt. (Papanikolaou *et al.* 1988b, Ganas *et al.* 2004); 4) the 14 km long Afidnai fault, which bounds the Afidnai plain to the south and the Athens Plain to the north (see also Ganas *et al.* 2005).

All these faults have produced a cumulative vertical throw of about 2 km (Fig. 3). The Alpine formations at the top of the Parnitha Mt at about 1400 m are buried several hundred meters beneath the present-day sea level and below the Quaternary and the Neogene sediments near the coast of Oropos. The footwall of the Milesi fault is the northernmost Alpine outcrop observed (Figs 1, 3). These active faults are widely distributed and sharing the strain so that they all exhibit relatively low slip-rates (<0.5 mm/yr). However, they are competitive structures distributed across strike lying on each other's stress shadows (Cowie 1998), implying that in the future one of them will prevail.

It is important that the above set of E-W trending active normal faults in northern Attica is constrained exclusively within the non-metamorphic Alpine units. It is also interesting that these faults taper out to the east as approaching the NNE-SSW trending tectonic contact that separates them from the metamorphic units. These observations are in agreement with the disruption model of the Hellenic arc (see Papanikolaou and Royden 2007) which considers: i) the southern Evia / Northern Attica detachment as the southern boundary of the Central Hellenic Shear Zone activated during Miocene to Early Pliocene and ii) the E-W faults developed to the northwest of the detachment, activated since Late Pliocene. Therefore, this detachment indirectly seems to have a significant effect on the neotectonic structure, influencing the geometry, style and intensity of deformation.

4. The Afidnai fault and the Northern edge of the Athens Plain

4.1. The Athens Plain

In the Late Miocene an E-W trending fault from Attiko Alsos to Kamatero divided the Athens Plain in two sub-basins. The northern part subsided forming a series of lakes throughout Late Miocene-Pliocene, whereas the southern part was tilted to the south with Alpine formations at the northern footwall zone (Attiko Alsos or Tourkovounia) mostly covered by shallow marine sediments in the southern coastal zone of Pireas-Trachones (Papanikolaou *et al.* 2004b). In Early? -Middle Pleistocene the Kifissos river was formed, cutting through the now inactive E-W fault which previously acted as a barrier. As a result, both subbasins were merged into one, producing what is today known as the Athens Plain (Papanikolaou *et al.* 2004b). Overall, since Middle Pleistocene the Athens Plain is very similar to the present-day feature. However, one important modification that occurred since Middle Pleistocene concerns the initiation of activity and/or the speeding up of the Afidnai fault. This fault constrained the northern edge of the Athens plain as well as its drainage basin. Due to the backtilting, most of the footwall area is drained by an extensive channel network that flows southwards down the dip footwall slope, forming the upper part of the Kifissos network. The asymmetry of Kifissos drainage basin, where most channels flow

from Partnitha Mt to the southeast (Fig. 2), illustrates nicely this recent tilt. Also in Middle Pleistocene thick fans and terraces were formed. These planation surfaces observed up to the area south of Krioneri are offset by the Afidnai fault (Fig. 5). As soon as the E-W Kamatero fault became inactive in Middle Pleistocene, the Afidnai fault initiated or took over the extra strain increasing its rate of slip.

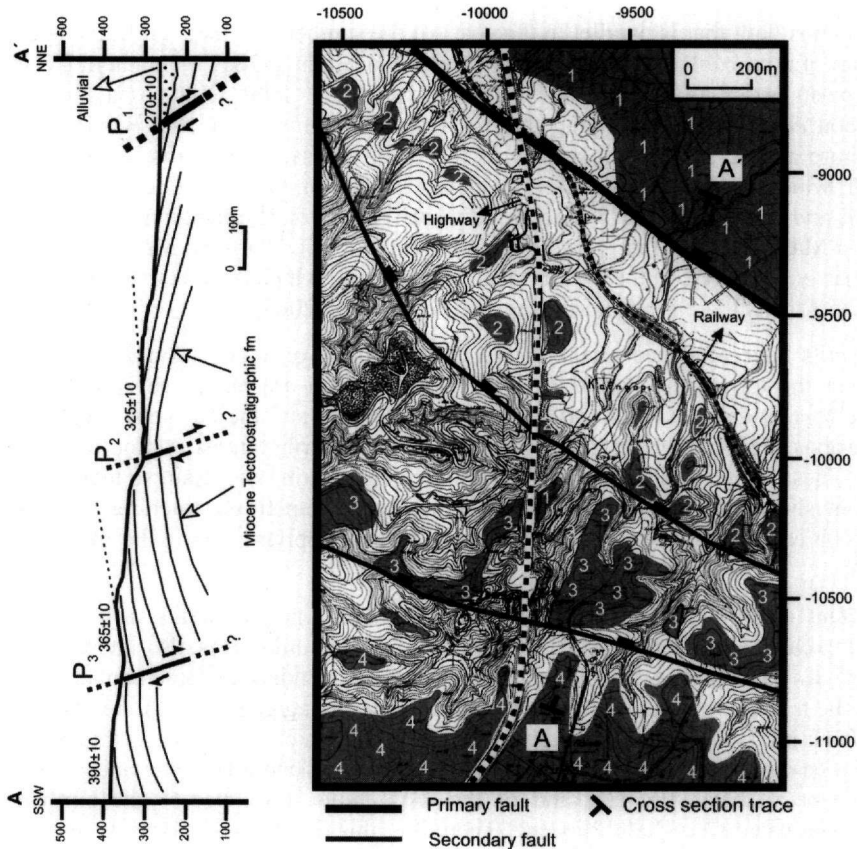


Figure 5 - 1:5000 topographic map (6436/3) and cross-section (A-A') towards the eastern part of the Afidnai fault, showing the traces of the major and the secondary faults and the offset planation surface (for location see Fig. 2)

4.2. The Afidnai fault

The Afidnai fault is a E-W trending normal fault, which downthrows to the north and exhibits all the characteristics of an active structure. This fault does generate an impressive morphological footprint, is oriented perpendicular to the present regional extension direction and bounds the Afidnai Quaternary basin. Even though no post-glacial scarps have been identified, there are signs in the geomorphology and drainage, attesting to its recent activity. In particular, the Charadros river is clearly deflected into a fault parallel flow direction due to the subsidence within the Afidnai hangingwall (Fig. 2). Moreover, the Charadros drainage divide is located only several hundred meters southwards from the Afidnai fault in its footwall, but several kilometers northwards from the fault trace in its hangingwall (Fig. 2). The proximity of the footwall watershed to the fault trace, is due to the tilted footwall that produced southward flow directions, which drain the footwall away from the hangingwall. In conclusion, it is suggested that the fault is active for several reasons such as: i) its geometry (~E-W trending fault parallel to all other major regional active faults, see Mariolakos and Papanikolaou 1981, 1987 and compatible with the

regional ~N-S stress field, Ambraseys and Jackson 1990, Roberts and Ganas 2000, Ganas *et al.* 2004), ii) its morphological expression (Fig. 6), iii) the fact that it bounds Quaternary sediments, and iv) its drainage pattern.

The Afidnai plain is a local base level. The short distance of the drainage divide on the footwall of the Afidnai fault from the fault trace shows that the dominant route for sediments to enter the adjacent basin in the hangingwall is via short, steep footwall catchments. Therefore, the volume of sediment supply into the Afidnai basin is restricted, resulting into a basin of limited extent and thickness. On the other hand, the limited thickness and spatial distribution of the Quaternary sediments also imply that this base level is recently established. The latter indicates that this fault is relatively young (Early? -Middle Pleistocene fault) that is also in agreement with the paleogeography of the Athens Plain.

The rate at which a fault slips fundamentally determines the seismic hazard because average recurrence intervals tend to decrease as slip-rates increase (e.g. Cowie and Roberts 2001). Unfortunately, the lack of characteristic stratigraphic horizons and postglacial scarps limits our ability to calculate accurately the finite throw or the throw-rate. However, the study of geomorphic features, such as terraces/planation surfaces, can provide us with a first order pattern of displacement. The dense vegetation cover of the Parnitha national park and the highly disturbed topography, due to the existing old clay quarries west of the Athens-Thessaloniki highway, severely limit our observation points. Therefore, we focused our terrace mapping mostly east of the highway, where geomorphology is still relatively undisturbed, even though we are aware that we are closer to the eastern tip of the fault, rather than to the center of it. Detailed 1:5000 mapping of the planation surfaces on the northern edge of the Athens Plain revealed the presence of the main fault at the base of the hillside and two secondary faults on the footwall of the main fault. It is unclear whether the secondary faults are still active or that activity may have shifted to the lower major fault following a progressive hangingwall directed migration within the fault zone (Stewart and Hancock 1994). The Afidnai fault separates the major Athens plain from the minor Afidnai-Lake-Marathon Plain. The present day Afidnai basin on the immediate hangingwall of the Afidnai fault lies at 270 ± 10 m, whereas the northern edge of the Athens basin is at 390 ± 10 m, implying a minimum finite throw of 120 m. Half of this throw is accommodated by the main fault at the base and the other half from the two secondary faults towards the footwall, which are probably linked at depth with the main fault. Considering an Early Pleistocene age (between 1.0 to 1.5 Ma) of the planation surface (see also 4.1), we calculate a long-term throw-rate of 0.08-0.12 mm/yr. However, this rate is extracted from a locality, which is closer to the eastern tip of the fault, rather than to the center of it (Fig. 6). The latter suggests that we may underestimate the rate towards the center of the fault (Papanikolaou and Roberts 2006), because fault slip and slip-rates vary along strike the fault and increase systematically along strike from zero at the fault tips, to maxima towards its center (e.g. Cowie and Scholz 1992). The latter can be nicely observed on the panoramic view of the fault taken from the area of Kapandriti, showing how displacement progressively decreases to minima as approaching the eastern tip of the fault (Fig. 6). Considering that the finite throw towards the center of the fault is about 3 times higher (~360 m, see also Ganas *et al.* 2005 who calculate a 350 m relief based on DEM analysis), the maximum rate is about 0.30 mm/yr. Following the above discussion, we estimate a mean rate of 0.15-0.20 mm/yr for this fault. This value is similar to the value extracted from Ganas *et al.* (2005). They based on DEM morphometry and made an empirical suggestion that slope angle is a function of the long-term fault slip-rate, which varies between 0.13-0.30 mm/yr. For the Afidnai fault they calculated a 0.15-0.19 mm/yr, but using a conservative (4-5 Ma) age constrain. In such a low throw-rate, no postglacial fault scarps or well-shaped bedrock scarps with a free face can be revealed. This occurs because erosion and sedimentation rates are higher and outpace the fault throw-rate.

Even though this fault offsets the Miocene tectonostratigraphic formation, it seems that its eastern tip is located close to the detachment trace, showing once again that the detachment plays a fundamental role in the region (Fig. 1). Further to the east and north from the Dionysos fault, two

NW-SE trending possibly active faults may take over the strain from the Afidnai fault, but no straightforward link to the Afidnai fault can be supported. They leave a rather subtle trace in the morphology and no significant throws can be calculated.



Figure 6 - Panoramic view of the Afidnai fault geometry (photo taken from Kapandriti). Footwall elevation tends to decrease away from the fault centre. Therefore, displacement decreases to minima as approaching the eastern tip of the fault

4.3. Seismic hazard assessment

Following the historical catalogue, no large events ($M > 6.5$) or severe damages have been recorded in NE Attica (Galanopoulos 1955, Papazachos and Papazachou 1997). Only two events have been recorded, the Oropos event in 1938 and another earthquake in 1705. The 1938 Oropos event ($M_s = 6.0$) probably ruptured part of the Oropos offshore fault or the shorter Milesi fault (Fig. 1), generating macroseismic intensity VIII and 18 fatalities (Papazachos and Papazachou 1997). No primary surface ruptures have been observed or documented, but several landslides occurred particularly in the area of Malakasa. As far as the 1705 event (Ambraseys and Jackson 1997) is concerned, there is uncertainty concerning both the magnitude ($M \sim 6.4$ according to Papazachos and Papazachou 1997) and predominantly the epicentral area. Minor damages were inflicted both to the towns of Athens and Chalkida, so that researchers place the epicenter towards the NE flanks of the Parthitha mountain. It is possible that the Afidnai fault could have ruptured during the 1705 event, however the limited data extracted from the historical record cannot confirm or deny such an interpretation (e.g. it could be also the Malakasa-Avlona fault). The historical record is considered complete for shallow events ($h < 60$ km) for $M > 6.5$ since 1845 and for $M > 7.3$ since 1500 (Papazachos *et al.* 2000). Therefore, it is also possible that an earthquake of $M = 6.0$ to 6.4 could have occurred before the year 1845 and for several reasons has not been recorded.

The worse case scenario implies the rupture of the entire 14 km long Afidnai fault. Based on the equation between Magnitude (M) and surface rupture length (SRL) of the worldwide dataset of Wells and Coppersmith (1994), this fault can generate a $M = 6.4$ event. A similar magnitude ($M_s = 6.5$) is also calculated based on Pavlides and Caputo (2004) empirical relationships from normal faulting events of the Aegean, which may be more representative of the local geotectonic conditions. Ganas *et al.* (2005) estimated the fault as 11 km and calculated a similar Magnitude ($M_s = 6.4$). Considering a mean slip-rate in the order of 0.15-0.20 mm/yr and that all slip is released in earthquakes of the maximum magnitude ($M = 6.4$), therefore assuming a worst-case scenario, we calculate a recurrence interval of about 2000 years.

Taking into account that the completeness of the historical record for such magnitude events covers time periods of a few hundreds years and probably shorter than 500 years (e.g. Papazachos *et al.* 2000), it is evident that the recurrence interval is only a fraction of the completeness period.

Therefore, either this fault has not been activated in historical times or it was activated, but for whatever reason such an event was not recorded in the historical catalogues. Finally, if this fault indeed ruptured in 1705, then the probability for a future event in the near future generated by this source is rather negligible.

5. Discussion-Conclusions

Relief in NE Attica is mainly influenced by fault geometry and fault spacing. Drainage basins are highly asymmetric, producing a combination of fault parallel (E-W) and fault perpendicular (N-S) flow directions. The closely spaced faults produce secondary narrow highs and lows due to the interference between displacement fields, influencing also the drainage divide. The Afidnai fault bounds the northern edge of the Athens Plain and separates it from the secondary Afidnai Plain. It is active, but exhibits a low slip-rate. We have used offset Pleistocene terraces to calculate a long-term throw-rate of 0.1 to 0.3 mm/yr for the Afidnai fault, depending on the actual terrace Pleistocene age estimates and our location along strike the fault. This fault is probably active since Early? -Middle Pleistocene and based on a worst-case scenario, it can generate events of the maximum magnitude ($M=6.4$), about every 2000 years. Therefore, the completeness of the historical record is much shorter than the mean earthquake recurrence interval of the Afidnai fault.

The NE Attica is also divided by a major NNE-SSW detachment fault that separates the metamorphic units to the east from the unmetamorphic units to the west. This detachment also separates the E-W trending faults towards the western part from the NW-SE trending less active faults towards the eastern part. It was active in Late Miocene-Early Pliocene and produced several hundred meters of debris-flow deposits. It is important that the set of E-W trending active normal faults in northern Attica is constrained exclusively within the non-metamorphic Alpine units and that these faults taper out to the east as approaching the detachment zone. Therefore, this detachment indirectly seems to have a significant effect on the neotectonic structure, influencing the geometry, style and intensity of deformation.

Finally, it is interesting to note that this NNE-SSW trending detachment influences also the seismicity pattern. More precisely, it coincides with the line separating zone I (lowest category of seismic risk) from zone II (intermediate zone) of the national seismic building code (EAK-2003), which have been compiled based on the seismicity level (E.P.P.O.-A.C.E.G., 2001). This is also in agreement with the damage pattern of the 1999 Athens Earthquake, where this detachment further to the south (e.g. inside the Athens Plain) played an important role to the intensity distribution (Papanikolaou *et al.* 1999, Marinos *et al.* 1999), forming a boundary between higher and lower intensities. More precisely, significantly higher damages were produced to sites located both along strike its trace and west from it, within the non-metamorphic units, compared to the sites situated east of the detachment. Therefore, this detachment even though it is inactive, still exerts a significant influence on the present day geological, geomorphological, neotectonic and seismicity pattern, showing how old and new structures interrelate and interact.

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7. References

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