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# RELATIONSHIPS BETWEEN PETROGRAPHIC AND PHYSICOMECHANICAL PROPERTIES OF BASIC IGNEOUS ROCKS FROM THE PINDOS OPHIOLITIC COMPLEX, NW GREECE

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#### Abstract

The relationships between petrographic and physicomechanical properties of basic igneous rocks from the Pindos ophiolitic complex, northwestern Greece, were investigated. The mineralogical composition was quantified from one polished thin section per sample with a polarizing microscope, by using the point count method. Textural description was also carried out by using both polarizing and scanning electron microscope. The same samples were also tested to determine moisture content, water absorption, specific gravity, total porosity, void ratio, uniaxial compressive strength, Los Angeles, micro-Deval, PSV and Sand Equivalent values. The relationships between these properties and the petrographic characteristics are described by simple regression analyses. The results indicate that plagioclase, chlorite, quartz and actinolite contents of the studied rock-types significantly influence their physicomechanical properties. Additionally, some textural parameters such as the mean grain size and the complexity of grain boundaries influence their mechanical strength.

Key words: aggregates, quantitative petrography, dolerites, troctolites.

#### Περίληψη

Διερευνήθηκαν οι σχέσεις μεταξύ πετρογραφικών χαρακτηριστικών και φυσικομηχανικών ιδιοτήτων βασικών πετρωμάτων από το οφιολιθικό σύμπλεγμα της Πίνδου, στη βορειοδυτική Ελλάδα. Η ορυκτολογική σύσταση ποσοτικοποιήθηκε από μία λεπτή τομή για κάθε δείγμα, με τη χρήση πολωτικού μικροσκοπίου και με βάση τη μέθοδο "point count". Επίσης πραγματοποιήθηκε ιστολογική περιγραφή με τη χρήση τόσο πολωτικού όσο και ηλεκτρονικού μικροσκοπίου σάρωσης. Στα ίδια δείγματα προσδιορίστηκαν η περιεχόμενη υγρασία, η υδαταπορροφητικότητα, το ειδικό βάρος, το ολικό πορώδες, ο λόγος κενών, η αντοχή σε μοναζονική θλίψη και οι τιμές Los Angeles, micro-Deval, PSV και ισοδύναμου άμμου. Οι σχέσεις μεταξύ των παραπάνω ιδιοτήτων και των πετρογραφικών χαρακτηριστικών προσδιορίστηκαν με απλή ανάλυση παλινδρόμησης. Τα αποτελέσματα υποδεικνύουν ότι η περιεκτικότητα σε πλαγιόκλαστο, χλωρίτη, χαλαζία και ακτινόλιθο των υπό μελέτη λιθοτύπων, επηρεάζει αζιοσημείωτα τις φυσικομηχανικές τους ιδιότητες. Επιπλέον, η μηχανική τους αντοχή επηρεάζεται από ορισμένες ιστολογικές παραμέτρους, όπως το μέσο μέγεθος κόκκων και την περιπλοκότητα των μεταξύ των κόκκων ορίων.

Λέζεις κλειδιά: αδρανή υλικά, ποσοτική πετρογραφία, δολερίτες, τροκτόλιθοι.

#### 1. Introduction

The physicomechanical properties of rocks are the most important parameters in the design of ground workings and in the classification of rocks for engineering purposes. They are dependant on the mineralogy, texture (size, shape and arrangement of mineral grains, nature of grain to grain contacts, degree of grain interlocking), alteration and deformation of the source rock (Shakoor and Bonelli 1991, Haney and Shakoor 1994, Tuğrul and Zarif 1999, Smith and Collis 2001, Miskovsky *et al.* 2004, Tsikouras *et al.* 2005, Al-Oraimi *et al.* 2006, Rigopoulos *et al.* 2006). The correlation between the physicomechanical properties and quantitative petrography of rocks is dependant on various petrographic properties simultaneously. Thus, most studies have mainly concentrated on groups of rock or geological property: Irfan and Dearman (1978) on weathering, Brattli (1992) on basic igneous rocks, Shea and Kronenberg (1993) on schists and gneisses, Tuğrul and Zarif (1999) on granitic rocks, Åkesson *et al.* (2003) on foliation of granitic rocks and Räisänen (2004) on hybridisation and hornblende granites.

The aim of this study is to determine the relationships between the petrographic features and the physicomechanical properties of the samples studied, which include 4 dolerites and 2 troctolites. Physicomechanical tests and traditional petrographic description of thin sections with a polarizing microscope were performed. Point counting analyses were carried out for the quantification of the mineralogical data. In addition, scanning electron microscopy (SEM) studies were conducted, in order to determine mineral chemistry. Besides, the surface roughness and porosity of the rocks were observed by using SEM images.

### 2. Geological setting

The investigated samples comprise basic members of the Pindos ophiolite, one of the best preserved ophiolite sequences of Greece, which is located in the Subpelagonian Zone. The Pindos ophiolite complex consists of three main tectonic units: the Dramala Complex, the Aspropotamos Complex and the Loumnitsa Unit that are structurally underlain by the Avdella mélange (Jones and Robertson 1991). The Aspropotamos Complex is distinguished in eastern and western Aspropotamos Complex (Kostopoulos 1989).

The Dramala Complex consisting of harzburgite peridotite with abundant podiform dunite and some plagioclase lherzolite, is volumetrically the predominant lithology in the Pindos ophiolite complex (Jones and Robertson 1991).

The eastern Aspropotamos Complex (Kostopoulos 1989) consists of a sequence of mafic and ultramafic cumulates overlain by gabbros and minor plagiogranites with sheeted dykes, capped by a series of extrusive rocks. The western Aspropotamos Complex displays a dismembered character, including olivine-bearing isotropic gabbros with plagiogranite pods. These are stratigraphically overlain by sheeted dykes that pass upward into pillow lavas.

The Loumnitsa Unit comprises the metamorphic sole at the base of Dramala and Aspropotamos Complexes, consisting of lower amphibolite and greenschist phases metamorphic rocks of igneous and sedimentary origin. Radiometric dating on hornblende separates from the amphibolite, suggest a Late Jurassic age (165±3 Ma; Spray *et al.* 1984).

This study is concentrated on Spileo and Koridallos areas, which belong to the eastern Aspropotamos Complex and to the Avdella mélange respectively. Extensive fieldwork in these areas focused mainly on the distribution and mode of development of the mafic lithologies, as well as their relation to the adjacent formations, and resulted in detailed geological maps (Fig. 1). Sampling was performed north of the Veneticos valley (W of Spileo village) in a well preserved sheeted dyke complex (Fig. 1a) and North of the Koridallos village in troctolites (Fig. 1b). In the Veneticos valley, the sheeted dyke complex was further divided in three subunits, which are shown schematically in the geological map. Their boundaries are gradational and compose a

sequence with fine-grained dolerite and pillow lavas, a fine- to medium-grained dolerite and a medium-grained dolerite and gabbro. The dolerite dykes show well preserved chilled margins and are locally up to 3m wide. Two sets of dykes are present: an older one trending NW-SE is intruded by a younger set in a NNW-SSE direction. Intense deformation is imprinted by 1-5 m wide cataclastic zones striking NW-SE and N-S to W-E trending joints.



Figure 1 – Geological – sampling map of the a) Spileo and b) Koridallos areas from the Pindos ophiolite complex

The troctolite exposure of the Koridallos area is mapped as a discrete tectonic slice ( $\sim 1 \text{ km}^2$ ) into the Avdella mélange of the Pindos ophiolite. In well preserved outcrops, troctolite shows layering. The layering is dominantly marked by distinctive alternations of plagioclase- and olivine-rich layers. This tectonic block is surrounded by a chaotic mixture of tectonically juxtaposed clasts of Triassic limestones, cherts and pillow lavas in a deformed mudstone-siltstone matrix which characterizes the Avdella mélange.

## 3. Petrography

The mineralogical and textural characteristics of the samples were studied by optical and scanning electron microscopy. The point count method under polarizing microscope was used to determine the modal composition. Approximately 300 equally distributed points were counted in each polished thin section. A summary of the mineralogical analysis results is given in Table 1.



Figure 2 – Polarized photomicrographs of samples a) PVN1a, b) PVN2a, c) PKD1, d) PKD2 (Pl: plagioclase, Ol: olivine, Act: actinolite, Qz: quartz) and Secondary Electron Images showing surface roughness of e) the dolerites (sample PVN1a) and f) troctolites (sample PKD1) from the Pindos ophiolitic complex

Dolerites from the area west of Spileo are fine- to medium-grained isotropic rocks with subophitic texture (Fig. 2a, b). Laths of subhedral plagioclase and interstitial anhedral clinopyroxene comprise the subophitic groundmass. The plagioclase crystals are moderately sericitized. Hydrothermal alteration resulted in the development of chlorite, quartz, epidote, actinolite, clinozoisite and secondary titanite. Frequently these alteration products (principally epidote and

quartz) fill veins. Small amounts of opaque minerals participate, too. Although alteration has taken place, the interlocking texture has been preserved. Micro-cracks are mainly intragranular and sparsely intergranular. Two groups of dolerites have been distinguished according to their secondary minerals contents. The first group (Fig. 2a) displays significant amounts of actinolite and quartz, while the second (Fig. 2b) is characterized by negligible amounts of actinolite and variable amounts of chlorite. In the first group, most of the clinopyroxene crystals are completely pseudomorphically replaced by actinolite.

Troctolites from the area north of Koridallos are medium- to coarse-grained rocks with cumulate texture. The only textural difference between the two troctolite samples is that PKD1 (Fig. 2c) has higher mean grain size in comparison to PKD2 (Fig. 2d). They are composed mainly of plagioclase and olivine. The cumulate texture is distinguished as alternations of plagioclase- and olivine-rich layers, however it is not always well developed. Grain boundaries are straight to slightly curved and micro-cracks are usually intragranular and rarely intergranular. Clinopyroxene and opaque minerals participate in minor concentrations. Serpentine, chlorite and quartz comprise alteration products. Olivine crystals exhibit various degrees of alteration to serpentine, while epidote and quartz fill veins through the rock.

The surface roughness and porosity of the samples studied were observed by using SEM images. Dolerites, which are fine- to medium-grained rocks, have higher surface roughness and porosity than troctolites, which are medium- to coarse-grained rocks (Figs 2e, f).

Table 1 – Mineralogical composition (in vol. %) of the investigated rock types (PI: plagio-
clase, Cpx: clinopyroxene, Act: actinolite, Ol: olivine, Chl: chlorite, Qz: quartz, Ep: epidote,
Serp: serpentine, Op: opaque minerals)

Sample	Rock	Pl	Срх	Act	Ol	Chl	Qz	Ер	Serp	Ор
PVN1a	Dolerite	47.7	13.7	23.3	-	4.3	7.7	1.3	-	2.0
PVN1b	Dolerite	49.3	15.3	21.0	-	3.0	7.0	1.7	-	2.7
PVN2a	Dolerite	50.7	38.0	0.3	-	7.3	1.0	0.7	-	2.0
PVN2b	Dolerite	46.7	36.0	2.0	-	10.3	0.7	2.0	-	2.3
PKD1	Troctolite	68.7	1.0	-	12.7	-	-	-	17.3	0.3
PKD2	Troctolite	71.7	0.7	-	10.7	-	-	-	16.7	0.3

# 4. Mineral chemistry

Major and minor element compositions of plagioclase, clinopyroxene, olivine, amphibole and chlorite crystals were analyzed on polished thin sections using a JEOL 6300 scanning electron microscope, equipped with EDS and WDS with a ZAF correction software information, at the Laboratory of Electron Microscopy and Microanalysis, University of Patras; operating conditions were accelerating voltage 15 kV and beam current 3.3 nA with 4  $\mu$ m diameter beam. Representative analyses are given in Table 2.

The plagioclase from the two groups of dolerites is labradorite-anorthite (An<sub>62-91</sub>), while there is a small proportion of secondary plagioclase of albite composition. The analyzed plagioclase crystals from the troctolites are bytownite-anorthite (An<sub>86-92</sub>). According to the Wo-En-Fs content and the I.M.A. classification system (Morimoto *et al.* 1988), the analyzed clinopyroxene crystals of the dolerites are augites, while the small proportion of clinopyroxene crystals of the troctolites are of diopside composition (not shown). The analyzed olivine crystals from the troctolite samples show a homogeneous composition from core to rim and are rich in forsterite component (Fo<sub>90-91</sub>). The analyzed amphiboles from the dolerites are of secondary origin and are classified as actinolite, according to the classification diagram of Leake *et al.* 1997 (not shown). On the classification diagram of chlorites (Hey 1954, not shown), the analyzed crystals from the dolerites plot as clinochlore.

Table 2 - Representative microanalyses of plagioclases, pyroxenes, olivines, amphiboles and chlorites from the dolerites and troctolites of the Pindos ophiolite complex (\*: below detection limit)

$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	1	14	Plagioclase	1				Pyroxene		1		Olivine			Amphibole	100		Chlorite	
1         1		<b>PVN1a</b>	PVN2a	PKD1	PKD2		<b>PVN1a</b>	PVN2a	PKD1	PKD2		PKD1	PKD2		<b>PVN1a</b>	PVN2a		<b>PVN1a</b>	PVN2
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	2	51.01	53.43	45.87	45.32	SiO <sub>2</sub>	54.12	52.62	54.67	53.80	SiO <sub>2</sub>	42.18	41.86	SiO <sub>2</sub>	52.45	53.05	SiO <sub>2</sub>	30.14	31.5
0.         110         230         240         120         330         M40         150         430         130         540         411         440         153         440         153         153         150         153         150         153         153         150         153         150         153	2	*	*	*	*	TiO <sub>2</sub>	•	0.68	•	*	FeO	8.69	8.63	TiO1		*	TiO1	•	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	03	31.06	28.94	34.25	35.08	Al <sub>2</sub> O <sub>3</sub>	1.96	4.85	3.31	2.66	MnO	*	*	Al <sub>2</sub> O <sub>3</sub>	5.43	4.11	Al <sub>2</sub> O <sub>3</sub>	17.23	16.9
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		0.61	0.82	0.98	0.42	FeO	5.97	5.55	2.27	2.33	MgO	48.58	49.12	FeO	16.46	15.32	FeO	19.27	19.2
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	0	*	*	*	*	MnO	*	*	*	*	Ca0	*	*	MnO	•	•	MnO	0.26	
0         148         128         135         136         Col         211         2.047         218         238         NiO         11.9         11.8         Col           0         27         42         1.46         106         8.0         9.1	0	*	*	*	# 1 :: 4	MgO	16.38	15.7	16.58	15.94	Cr203	•	*	MgO	12.10	13.94	MgO	18.83	19.2
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	•	14.88	12.78	17.35	17.56	CaO	21.12	20.47	23.18	23.88	NiO	*	*	CaO	11.79	11.68	CaO	•	
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	0	2.78	4.22	1.46	1.06	Na <sub>2</sub> O	*	*	*	*	Total	99.45	19.66	Na <sub>2</sub> O	•	*	Na <sub>2</sub> O	*	
$ \begin{array}{ c c c c c c c c c c c c c c c c c c c$		*	*	*	*	K20	*	*	•	*		4 Oxygens		K20	•	•	K20	*	
Normania         Solutionia         Solutionia         India         India <td>-</td> <td>100.34</td> <td>100.19</td> <td>16.66</td> <td>99.44</td> <td>Cr203</td> <td>0.51</td> <td>*</td> <td>0.64</td> <td>0.77</td> <td>Si</td> <td>1.03</td> <td>1.02</td> <td>Cr203</td> <td>•</td> <td>•</td> <td>NiO</td> <td>•</td> <td></td>	-	100.34	100.19	16.66	99.44	Cr203	0.51	*	0.64	0.77	Si	1.03	1.02	Cr203	•	•	NiO	•	
21)         242         213         247         210         240         310 <td>1</td> <td></td> <td>8 Oxygens</td> <td></td> <td></td> <td>Total</td> <td>100.1</td> <td>99.87</td> <td>100.7</td> <td>99.38</td> <td></td> <td>1.03</td> <td>1.02</td> <td>Total</td> <td>98.23</td> <td>98.10</td> <td>Cr203</td> <td>*</td> <td></td>	1		8 Oxygens			Total	100.1	99.87	100.7	99.38		1.03	1.02	Total	98.23	98.10	Cr203	*	
		2.31	2.42	2.12	2.10			Oxygens			Fe2+	0.18	0.18		23 Oxygens		Total	85.73	86.9
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		1.66	1.54	1.86	1.91	Si	1.98	1.92	1.96	1.96	Mn	i. N	ti e ta	Si	7.52	7.54		28 Oxygen:	
1.0 $2.00$ <		0.02	0.03	0.04	0.02	AIIV	0.02	0.08	0.04	0.04	Mg	1.77	1.78	AIIV	0.48	0.46	Si	6.24	6.
400         390         401         400         301         401         0.05 $1^{11}$ 0.06         0.12         0.10         0.08         Cr         1		-	1	2	, k	1	2.00	2.00	2.00	2.00	Ca	1	29 29		8.00	8.00	AIIV	1.76	1.
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		4.00	3.99	4.01	4.02	AIVI	0.06	0.12	0.10	0.08	Cr	i Ç	4	AIVI	0.44	0.23	ata arti	8.00	8.0
$\cdot$ <td></td> <td>•</td> <td>130</td> <td>1</td> <td></td> <td>Fe<sup>3+</sup></td> <td>i</td> <td>•</td> <td>•</td> <td></td> <td>N</td> <td>•</td> <td>•</td> <td>μ</td> <td></td> <td>•</td> <td>AI<sup>VI</sup></td> <td>2.45</td> <td>2.4</td>		•	130	1		Fe <sup>3+</sup>	i	•	•		N	•	•	μ		•	AI <sup>VI</sup>	2.45	2.4
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$		ł			1	H		0.02	1	L		1.94	1.96	Fe <sup>3+</sup>	0.42	0.68	Ті		
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$		0.72	0.62	0.86	0.87	c	0.01	a,"	0.02	0.02	Fo	90.88	91.03	c	911	į.	Ç	- sta	
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $		0.24	0.37	0.13	0.10	Mg	0.89	0.85	0.88	0.87	Fa	9.12	8.97	Mg	2.59	2.95	Mg	5.82	5.8
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	1			1		$\mathrm{Fe}^{2+}$	0.03	0.01		0.04				Fe <sup>2+</sup>	1.56	1.14	Fe	3.34	3.2
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$		76.0	0.99	0.99	0.96	Mn		b	•	•				Mn	•	-	Ni	Ĩ.	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		0.00	0.00	0.00	00.00		1.00	1.00	1.00	1.00					5.00	5.00	Mn	0.05	
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$		25.27	37.40	13.20	9.85	Mg	1		0.00	ľ				Mg		() ; ·	Ca	1	
Mn         -         -         -         Mn         -         -         K           Ca         0.83         0.8         0.89         0.93         Ca         1.81         1.78         11.1           Na         -         -         -         -         -         -         -         -         -         -         -         11.4           Na         - <td< td=""><td>1</td><td>74.73</td><td>62.60</td><td>86.80</td><td>90.15</td><td>Fe<sup>2+</sup></td><td>0.15</td><td>0.16</td><td>0.07</td><td>0.04</td><td></td><td></td><td></td><td>Fe<sup>2+</sup></td><td></td><td>1</td><td>Na</td><td></td><td></td></td<>	1	74.73	62.60	86.80	90.15	Fe <sup>2+</sup>	0.15	0.16	0.07	0.04				Fe <sup>2+</sup>		1	Na		
Ca     0.83     0.8     0.89     0.93     0.93     0.93     Ca     1.81     1.78     11.14       Na     -     -     -     -     -     -     -     -       K     -     -     -     -     -     -     -     -       Na     -     -     -     -     -     -     -     -       K     -     -     -     -     -     -     -       Fa     46.92     48.33     48.00     46.30     Ma     -     -     -       Wa     -     -     -     -     -     -     -     -     -       Wa     43.48     48.30     49.90     -     0.00     0.00     0.00     0.00						Mn	1	( ()	,	4				Mn			К	1	
Na         -         -         -         Na         -						Ca	0.83	0.8	0.89	0.93				Ca	1.81	1.78		11.65	11.5
K         -         -         -         200         200         200         200         200         200         200         200         200         200         200         200         E00						Na	T	5	•					Na		•			
0.97         0.96         0.96         0.97         Ca         -         -           En         46.92         46.83         48.00         46.30         Na         -         -         -           Fs         9.59         9.29         3.70         3.80         K         -         -         -           Wo         43.48         48.30         49.90         9.90         9.00         0.00         0.00						K	1			-					2.00	2.00			
En         46.92         46.83         48.00         46.30         Na         -							0.97	0.96	0.96	0.97				Са					
Fs         9.59         9.29         3.70         3.80         K         -         -           Wo         43.48         48.30         49.90         0.00         0.00         0.00						En	46.92	46.83	48.00	46.30				Na					
Wo 43.48 43.88 48.30 49.90 0.00 0.00						Fs	9.59	9.29	3.70	3.80				×					
						Wo	43.48	43.88	48.30	49.90					0.00	0.00			

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 Table 3 – Whole rock chemical analyses
 of representative samples of dolerites

 and troctolites from the Pindos ophio lite (\*: below detection limit)

<u> </u>	DETAIL	DETA	DIZDA
Sample	PVNIa	PVN2a	PKDI
Rock	Dolerite	Dolerite	Troctolite
Major El	emets (wt.%	6)	
SiO <sub>2</sub>	55.42	49.15	42.30
Al <sub>2</sub> O <sub>3</sub>	14.68	17.33	25.51
Fe <sub>2</sub> O <sub>3</sub>	9.33	8.46	2.89
MnO	0.13	0.12	0.03
MgO	6.70	7.56	9.63
CaO	8.30	10.69	11.88
Na <sub>2</sub> O	1.84	2.66	1.03
K <sub>2</sub> O	0.10	0.05	0.05
TiO <sub>2</sub>	0.20	0.98	0.01
$P_2O_5$	0.02	0.08	*
LOI	2.97	2.32	6.32
Total	99.69	99.40	99.65

### 5. Whole-rock chemistry

Whole-rock chemical analyses from 3 representative samples were performed by FUS-

ICP method (Table 3). According to their  $SiO_2$  content, the dolerites are typical basic rocks (49.15-55.42 %), however the troctolite exhibits lower  $SiO_2$  content (42.30 %). The higher  $Al_2O_3$  content of the troctolite relative to the dolerites is due to the higher plagioclase content of the former (see Table 1). Los on ignition ranges between 2.32 and 6.32 %, with the troctolites having the higher value due to the participation of serpentine after partial alteration of olivine.

#### 6. Physicomechanical properties

Knowledge of the physicomechanical properties of rock aggregates is essential in order to determine their quality. The results of the tests carried out are given in Table 4.

Physical properties investigated included moisture content (ASTM C 566), water absorption (ASTM C-128), bulk specific gravity (AASHTO T100-T85) and apparent specific gravity (ASTM C-127). Additionally, total porosity and void ratio were calculated. Three tests were performed for each property and the mean values were used. The troctolites exhibit lower moisture content, water absorption and total porosity and higher specific gravity values in comparison to the dolerites. Among the dolerite samples, PVN2a and PVN2b have higher total porosity and slightly higher water absorption and specific gravity values in comparison to PVN1a and PVN1b (Table 4).

Sample	PVN1a	PVN1b	PVN2a	PVN2b	PKD1	PKD2
Moisture content (%), n=3	0.52	0.44	0.51	0.56	0.31	0.22
Water absorption (%), n=3	0.53	0.45	0.55	0.59	0.34	0.23
Bulk spec. gravity (gr/cm <sup>3</sup> ), n=3	2.63	2.68	2.74	2.77	2.84	2.78
Ap. spec. gravity (gr/cm <sup>3</sup> ), n=3	2.55	2.62	2.62	2.64	2.79	2.72
Total porosity (%)	3.04	2.24	4.38	4.69	1.76	2.16
Void ratio	0.031	0.023	0.046	0.049	0.017	0.022
Los Angeles (%), n=1	10.72	10.55	11.37	12.42	12.84	12.50
micro-Deval (%), n=1	5.91	5.74	6.25	6.51	8.10	7.22
<i>P.S.V.</i> , <i>n</i> =2	57	56	55	56	43	46
Sand Equivalent (%), n=3	72	73	74	75	79	78
Un. Comp. Strength (MPa), n=6	155.59	162.78	158.34	149.56	160.07	187.02

Table 4 - Mean values of physicomechanical properties of the investigated rock types

The mechanical properties of the investigated rocks were determined by a variety of laboratory tests. Samples were first crushed with a laboratory jaw crusher. The material was then sieved to achieve the desired fraction and tested using the Los Angeles machine (ELOT EN 1097-02/1998), which defines an aggregate's resistance to fragmentation, the micro-Deval machine (ELOT EN 1097-01), which defines an aggregate's resistance to wear and the accelerated polishing machine (ELOT EN 1097-08), which gives a measure of an aggregate's resistance to the polishing action of vehicle tyres under conditions similar to those occurring on the surface of a road. Moreover, the sand equivalent values (ELOT EN 933.08), which define the clay-like fines in aggregates passing

ASTM Sieve No. 4 (4.75 mm) and the uniaxial compressive strength (ASTM D 2938-95) were determined. All of the samples studied have significantly low Los Angeles (10.55-12.84 %) and micro-Deval (5.74-8.10 %) values, indicating their high resistance to fragmentation and wear. However, the dolerite samples have lower Los Angeles, micro-Deval and Sand Equivalent and higher PSV values in comparison to the troctolites (Table 4).

### 7. Correlation analysis

Correlation analysis was used for the estimate of the influence of petrographic factors on the physicomechanical properties of the investigated ophiolitic rocks. Selected petrographic, physical and mechanical properties were plotted against each other in order to estimate one property from another. The correlation coefficients and best fit curves were calculated by the "least squares curves fit" method. The equations for the curves plotted and correlation coefficient values are given in the diagrams.

Figure 3a, b and c shows scatter plots for the relationships between the physical properties and the petrographic characteristics of the samples studied. Good positive correlations exist between the chlorite content and both the moisture content and water absorption values (Figs 3a, b). The porosity is an important factor in rock strength in that voids reduce the integrity of the material. It was found that an increasing content of quartz causes a decrease of total porosity (Fig. 3c).

The relationships between the mechanical properties and the petrographic characteristics have also been investigated. It was found that a rising content of plagioclase causes an increase in the Los Angeles and micro-Deval test values (Figs 3d, e), that means a decrease of the resistance of the rocks to wear and fragmentation. However, it should be mentioned that the coefficient of correlation ( $r^2$ ) between the plagioclase content and the Los Angeles test values is poor ( $r^2 = 0.47$ ). Inverse relationships exist between the quartz content and both the Los Angeles and micro-Deval test values (Figs 3f, g). In Figure 3h the actinolite content was plotted against PSV. It can be seen that the PSV values tend to increase with increasing actinolite content.

The various physicomechanical tests were also correlated with each other. Significant positive correlation was observed between the Los Angeles and micro-Deval test values (Fig. 4a). Besides, poor inverse relationship exists between the uniaxial compressive strength and total porosity (Fig. 4b).

#### 8. Discussion

Quantitative thin section studies can give a reliable forecast of rocks strength. The study presented has mainly been focused on the influence of the mineralogical composition of the investigated basic ophiolitic rocks on their physical and mechanical properties, by using the point count method. Additionally, textural description was carried out by using both polarizing and scanning electron microscope.

Two groups of dolerites were distinguished according to the amount of their secondary minerals, however their physicomechanical properties show only slight differences. The moisture content and water absorption tend to increase with increasing chlorite content (see Figs 3a, b) presumably due to absorption of water between the layers of chlorite. The negative correlation that is observed between the quartz content and the total porosity of the dolerite samples (Fig. 3c) is assigned to reduction of porosity by secondary quartz filling spaces between the other grains, analogous to what have been studied by Tuğrul and Zarif (1999).

The mechanical properties and especially the Los Angeles and micro-Deval values are influenced positively by the secondary quartz content (Figs 3f, g). This is due to the fact that quartz has no cleavage, so its abundance can result in higher strength. On the other hand, increasing plagioclase content influences the Los Angeles and micro-Deval values in a negative way (Figs 3d, e). These



Figure 3 - Relationships between petrographic and physicomechanical properties: chlorite content plotted against (a) moisture content and (b) water absorption, (c) quartz content plotted against porosity, plagioclase content plotted against (d) Los Angeles and (e) micro-Deval values, quartz content plotted against (f) Los Angeles and (g) micro-Deval values, (h) actinolite content plotted against PSV values (rhombs: group A of dolerites, triangles: group B of dolerites, squares: troctolites)



Figure 4 - Relationships between: (a) the micro-Deval and the Los Angeles values, (b) the porosity and the uniaxial compressive strength. Symbols same as in Figure 3

relationships are attributed to the perfect cleavage of plagioclase that apparently decreases the coherence of the rock. The dolerites have relatively lower plagioclase content (Table 1) and hence better mechanical properties (Table 4) compared to the troctolites. Similar relationships have also been investigated by Gunsallus and Kulhawy (1984) and Tuğrul and Zarif (1999). The replacement of clinopyroxene by actinolite in the dolerites, partly destroys their interlocking texture. This enables the rock particles to preserve their surface roughness, explaining the positive correlation between the PSV values and the actinolite content (Fig. 3h). The participation of actinolite and chlorite, together with the low mean grain size of the dolerites are the reasons for their higher PSV values compared to troctolites (Table 4).

Mean grain size is an important textural characteristic influencing the quality of aggregates. Most researchers have found that mechanical strength increases as mean grain size decreases (*e.g.*, Olsson 1974, Haraldsson 1984, Brattli 1992, Tuğrul and Zarif 1999, Räisänen 2004). The dolerites are fine- to medium- grained rocks and have slightly better mechanical properties, in comparison to the troctolites which are medium- to coarse-grained rocks (Table 4). Among the troctolites, PKD2 sample has lower mean grain size and slightly better mechanical properties in comparison to PKD1 (Table 4). Another feature observed in thin sections is the complexity of rocks grain boundaries. The lower Los Angeles and micro-Deval values of the dolerites compared to the troctolites is also attributed to the more complex grain boundaries of the former.

The relationship between anisotropy and the mechanical behaviour of aggregates has been studied by several researchers (*e.g.*, Schön 1996, Bohloli 2001, Åkesson *et al.* 2003). In this study, the mineral phases of the troctolites are not equally dispersed because of their cumulate texture. Grain boundaries between different phases result in better mechanical behavior of rocks relative to boundaries between the same phases. Based on this hypothesis, the cumulate texture of the troctolites may influence their strength, creating mechanically weak discontinuities.

#### 9. Conclusions

Relationships between the petrographic characteristics and the physicomechanical properties of selected basic igneous rocks from the Pindos ophiolite complex were determined by regression analysis. The samples studied include 4 dolerites and 2 troctolites. The results of this study can be summarized as follows:

 Mineralogical composition is one of the main factors controlling the rock physicomechanical properties. Abundance of secondary minerals influences the physical properties of the dolerites. The moisture content and water absorption tend to increase with increasing chlorite content and the total porosity tends to decrease with increasing quartz content.

- The mechanical properties of all samples studied are influenced negatively by the plagioclase content. Additionally, negative correlations were found between the secondary quartz content and both the Los Angeles and micro-Deval values of the dolerites. Significant positive correlation was also found between the actinolite content and the PSV values of the dolerites.
- The mean grain size influences the mechanical properties of the samples studied, with dolerites having higher strength than troctolites. Among the troctolites, PKD2 sample has lower mean grain size and slightly better mechanical properties in comparison to PKD1.
- The lower strength of the troctolites may also be a result of their cumulate texture and their less complex grain boundaries in comparison to the dolerites.

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